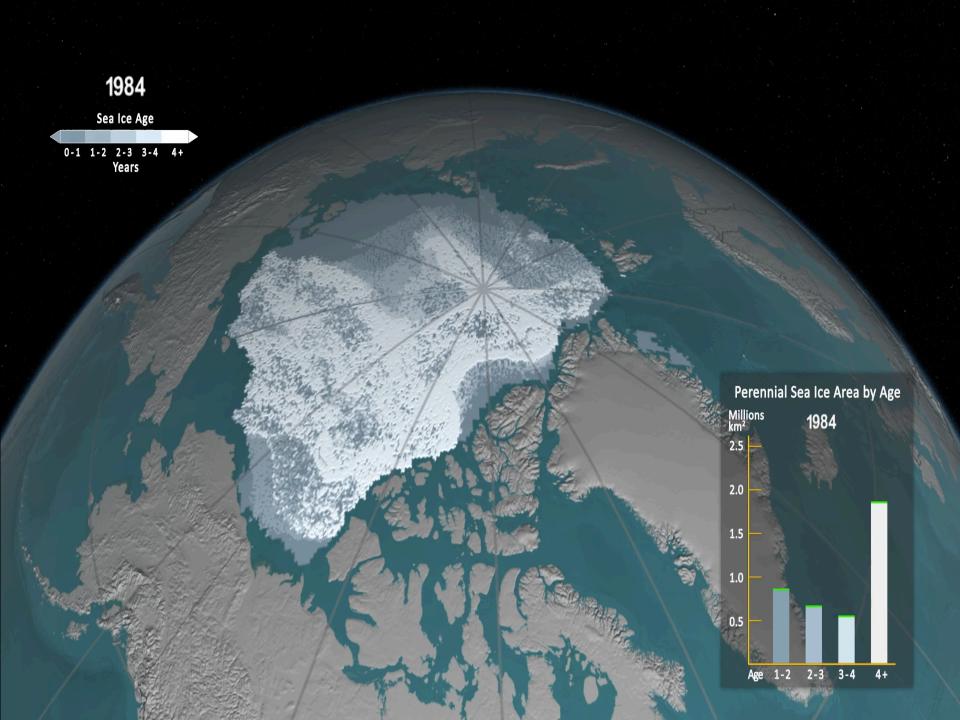


Arctic Low Cloud-Sea Ice-Stability Interactions and the Arctic Surface Energy Budget

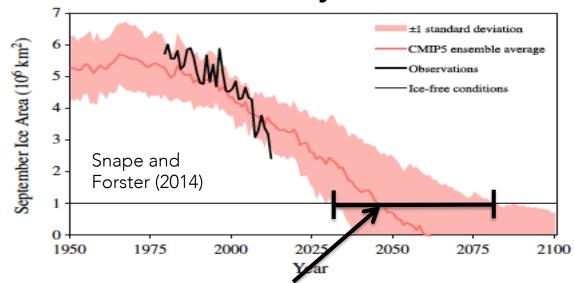
Dr. Patrick C. Taylor Science Directorate NASA Langley Research Center Hampton, VA May 17, 2017

patrick.c.taylor@nasa.gov

Acknowledgements: R. Boeke, B. Hegyi, S. Kato, K.-M. Xu and M. Cai



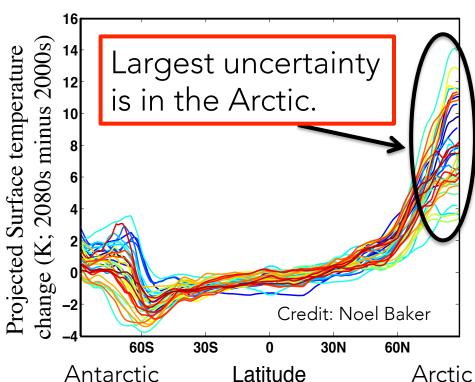
Oh, the uncertainty...



Projections of future Arctic sea ice decline and the timing of the first occurrence of a sea ice-free Arctic are very uncertain.

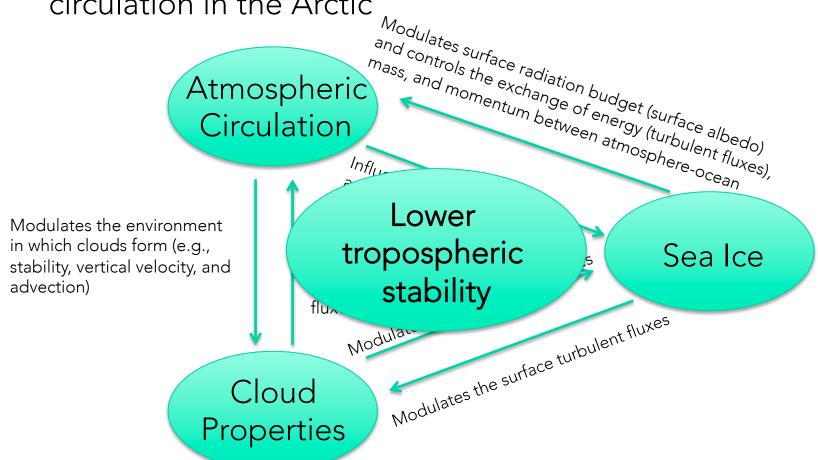
~50-year range in the projected first appearance of and ice-free Arctic

The large spread in climate model predictions of Arctic warming is attributed to model of sea ice melt and how it feeds back on the other components of the climate.



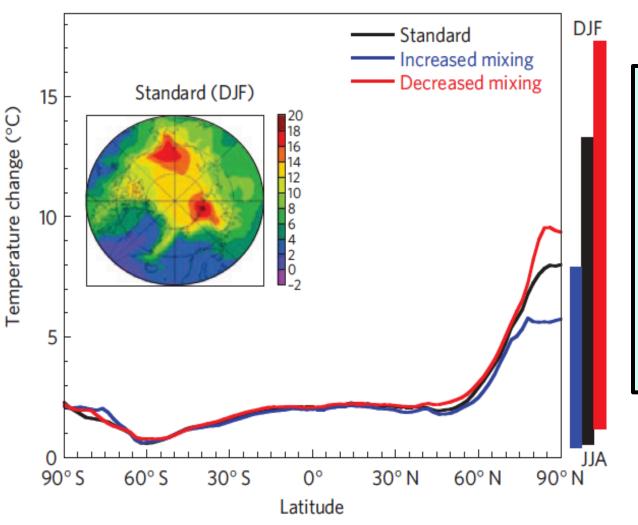
Another 3-legged stool?

Understanding the coupling between the cloud and circulation in the Arctic



In the Arctic, we cannot consider the interactions between clouds and the circulation without considering the sea ice state because sea ice influences both clouds and circulation.

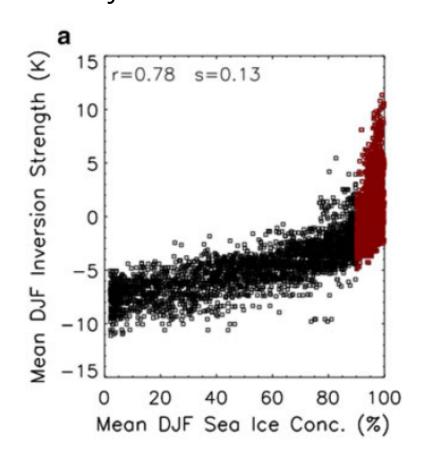
The impact of Lower Tropospheric Stability on Arctic Climate Change

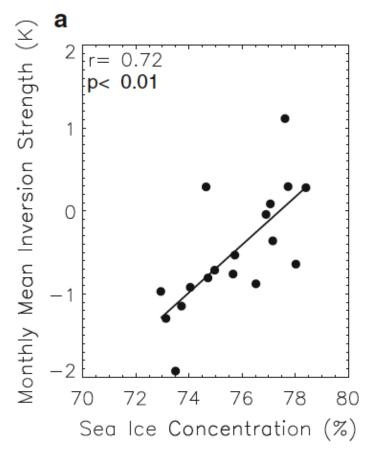


Stronger lower tropospheric stability and weaker turbulent mixing yields stronger Arctic warming.

Bintanja et al. 2011

Relationship between Lower Tropospheric Stability and Arctic sea ice cover

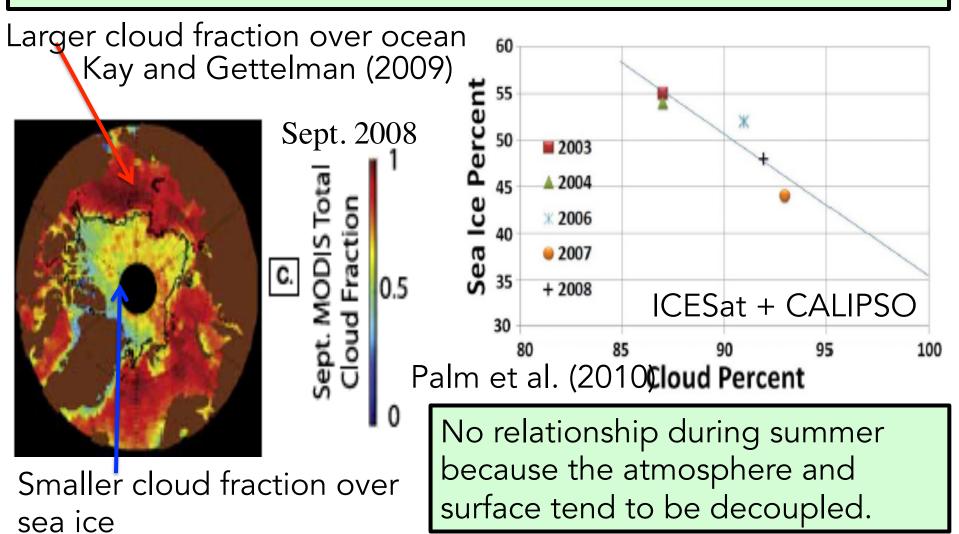




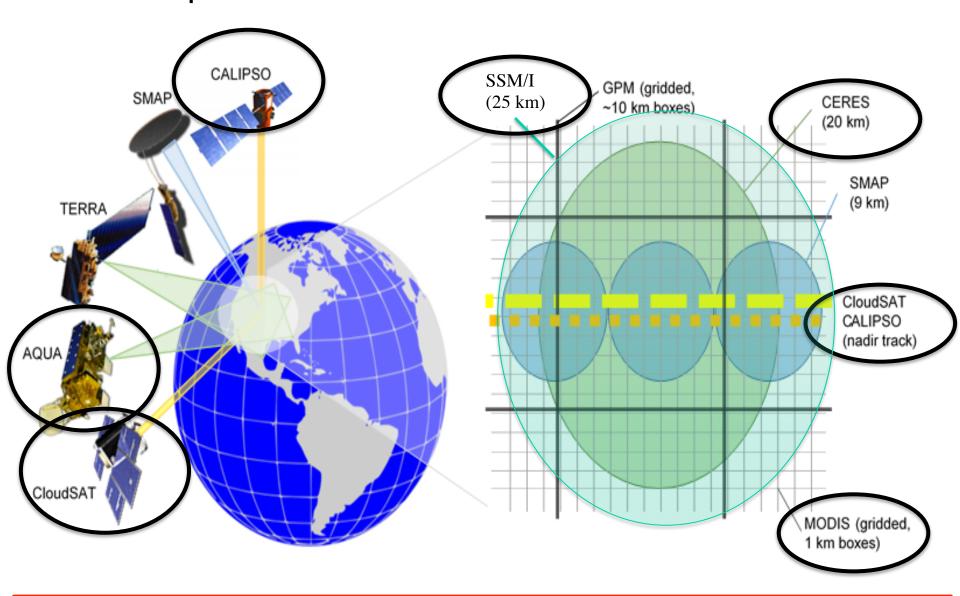
Lower tropospheric stability is strongly correlated in space and time the monthly average sea ice concentration (Pavelsky et al. 2011).

Sea ice-Cloud Interaction: Some Observational Evidence

Significant correlation between cloud fraction and the sea ice extent in AUTUMN: larger cloud fraction over open water and lower cloud fraction over ice.

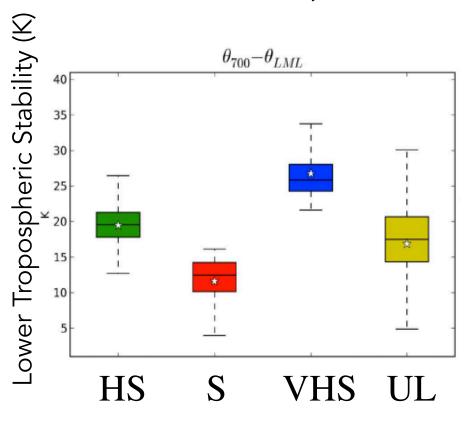


That's the power of...Data Fusion!?

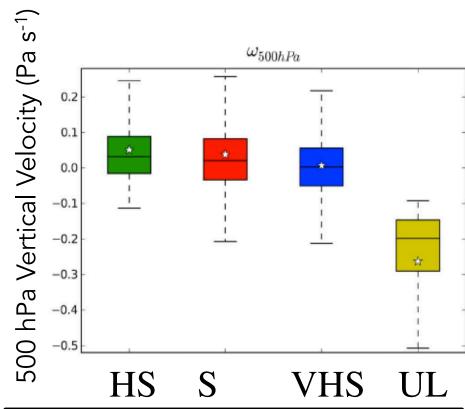


A new perspective is enabled by leveraging advances in data fusion made at NASA Langley Research Center—combining CALIPSO, CloudSAT, CERES, and MODIS.

Atmospheric Regimes (Barton et al. 2012)



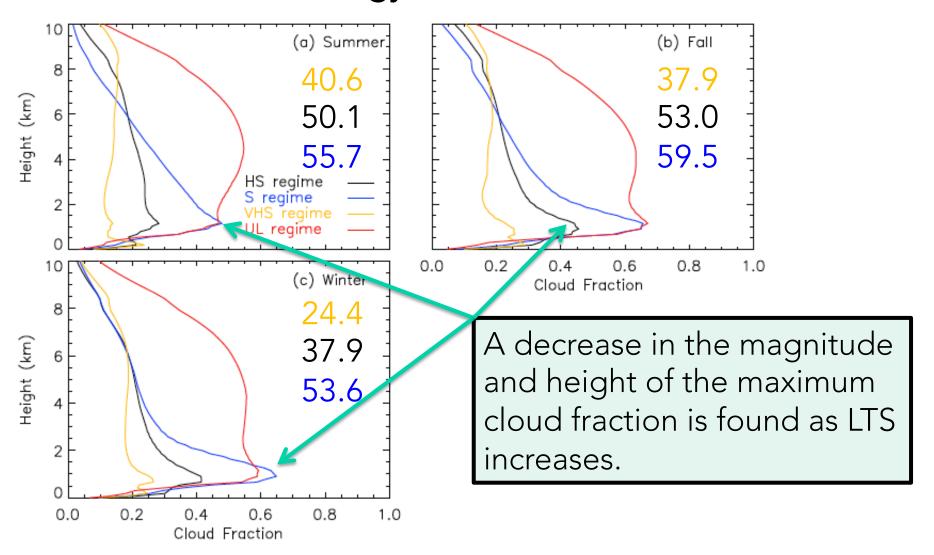
Atmospheric state regimes determined using K-means cluster analysis.



High Stability (HS): 16 K < LTS < 24 K Stable (S): LTS < 16 K Very High Stability (VHS): LTS > 24 K Uplift (UL): ω_{500} < -0.1 Pa s⁻¹

9

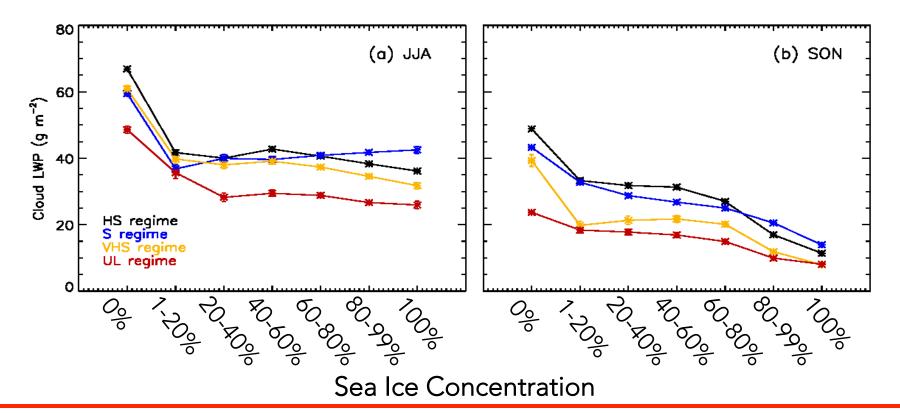
How does meteorology influence Arctic clouds?



Takeaway Message: Meteorology places a strong constraint on cloud behavior.

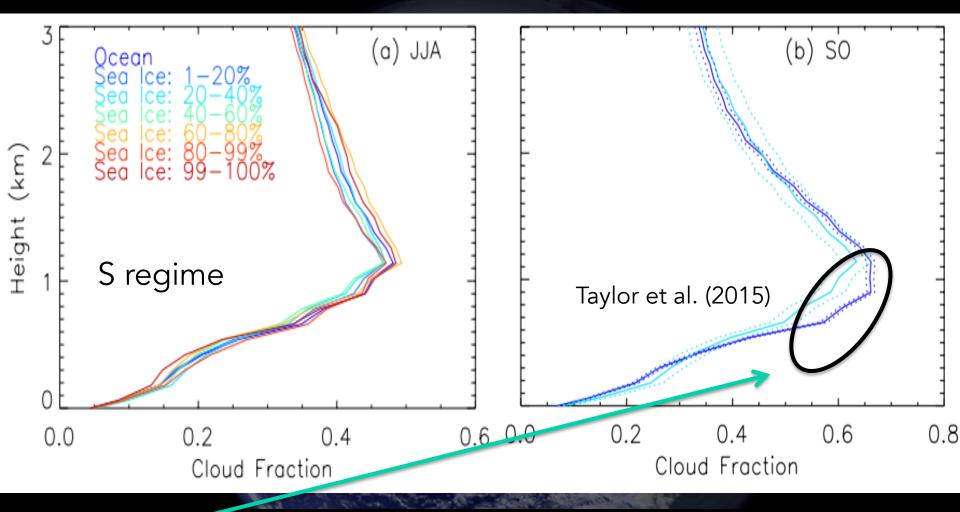
Compositing Methodology

- (1) Determine the Atmospheric Regime of each footprint using MERRA
- (2) Determine the instantaneous sea ice concentration from SSM/I retrieval
- (3) Average low cloud properties (cloud top < 3 km) within each atmosphere and sea ice concentration bin



The goal of the methodology is to retain as much process level information as possible by using satellite footprint level data, not monthly mean gridded.

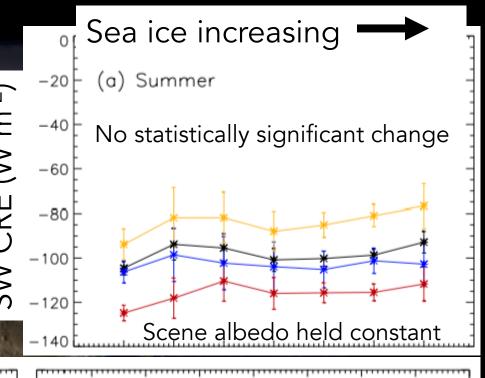
CF Vertical Profile vs. Sea ice

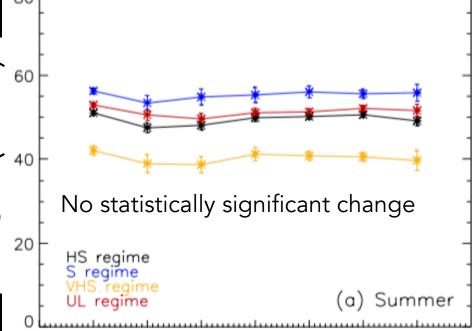


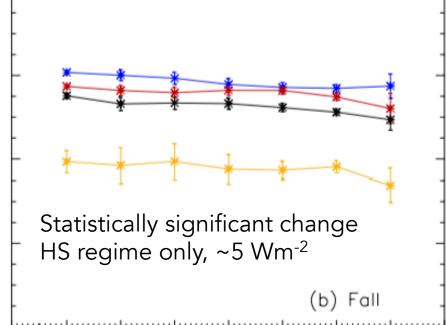
- General decrease in cloud fraction is found with increased sea ice concentration in in autumn, but no response in summer.
- Statistically significant differences at the 95% confidence interval are found at between 500 m and 1.2 km in autumn at 0% and 20-40% sea ice concentration.

CRE vs. SIC:

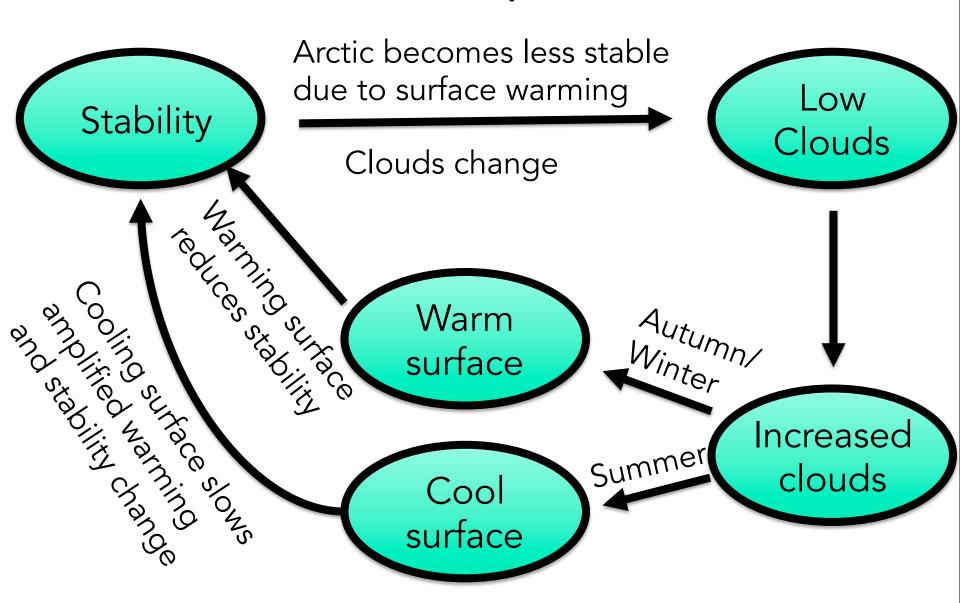
Large differences in LW and SW CRE in scene with different sea ice concentrations are found but few are statistically significant. Taylor (2017)







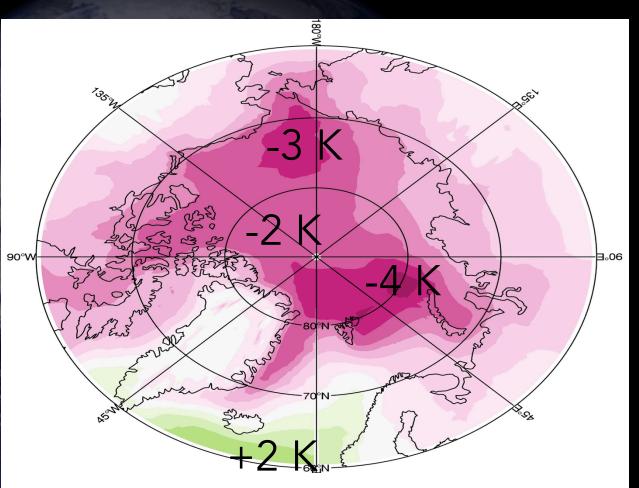
Stability-Cloud Response Pathway



Future changes in Arctic LTS:

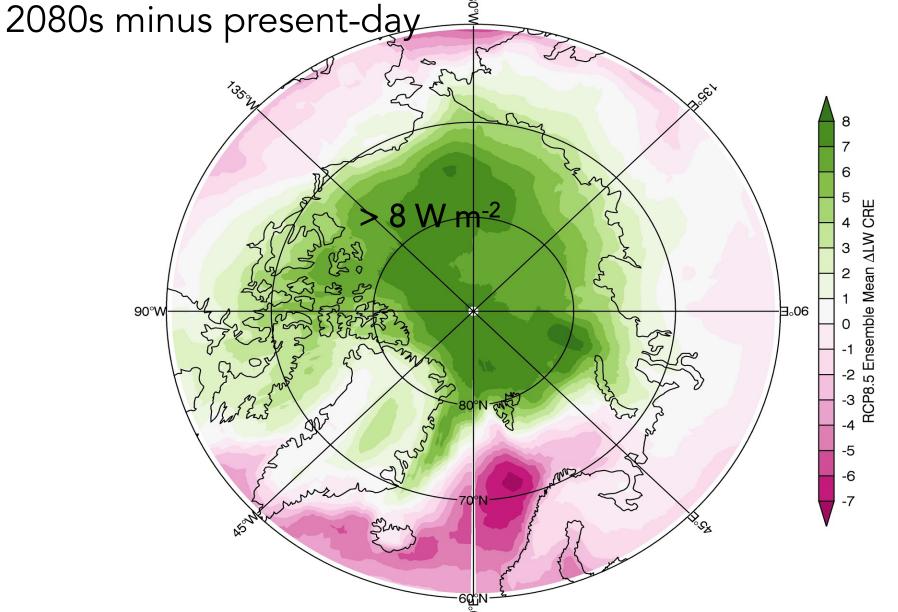
RCP 8.5

Overwhelmingly, CMIP5 models project significant decreases in LTS over the central Arctic Ocean.



Ensemble Average transient change in LTS

RCP8.5: Ensemble Average projected LW CRE change

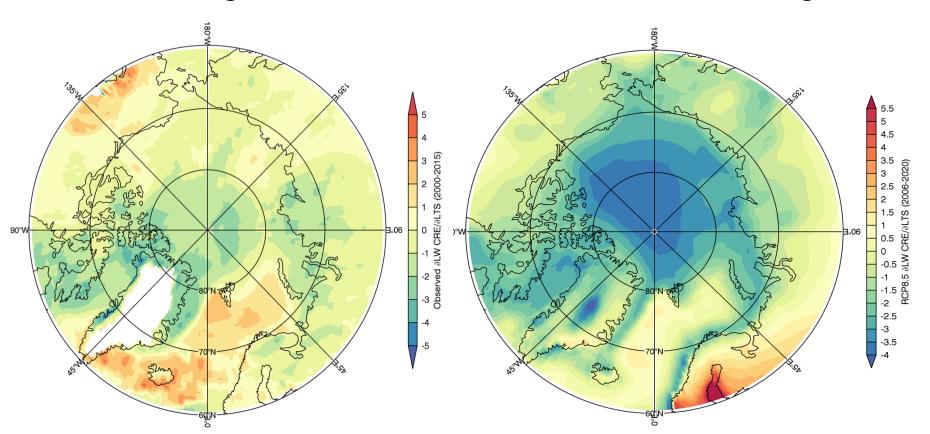


Largest changes in LWCRE are found in the central Arctic.

Inter-annual variability of LTS and LW CRE

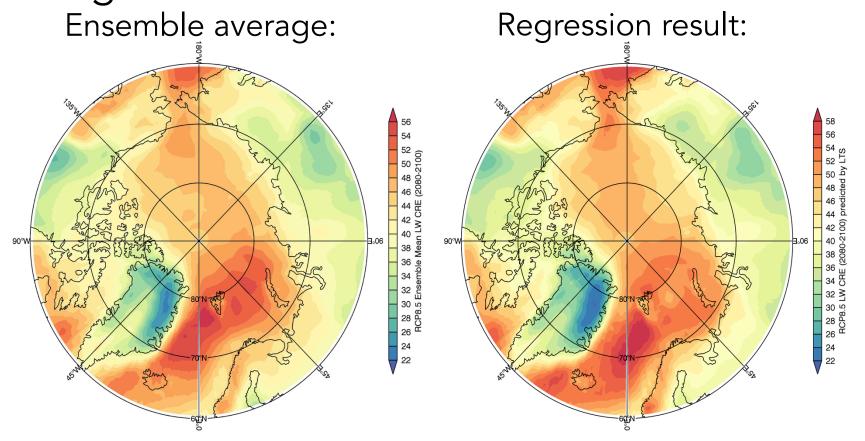
Observed regression

CMIP5 Ensemble Average



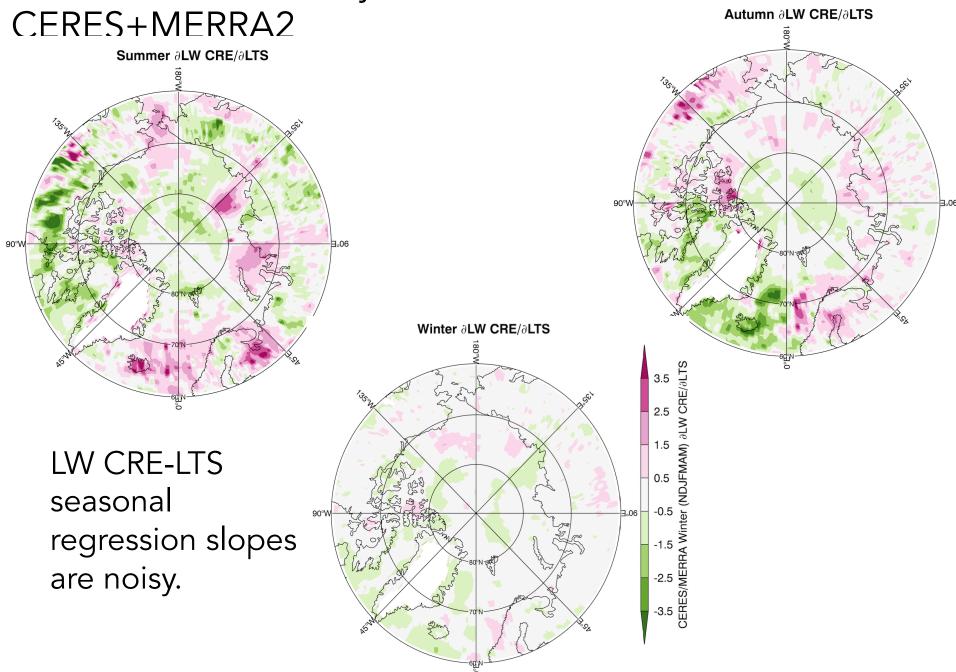
LW CRE-LTS regression slopes in climate models are more than 2 times larger than observations show over portions of the Arctic.

Can this regression be used to predict LW CRE change in 2080?

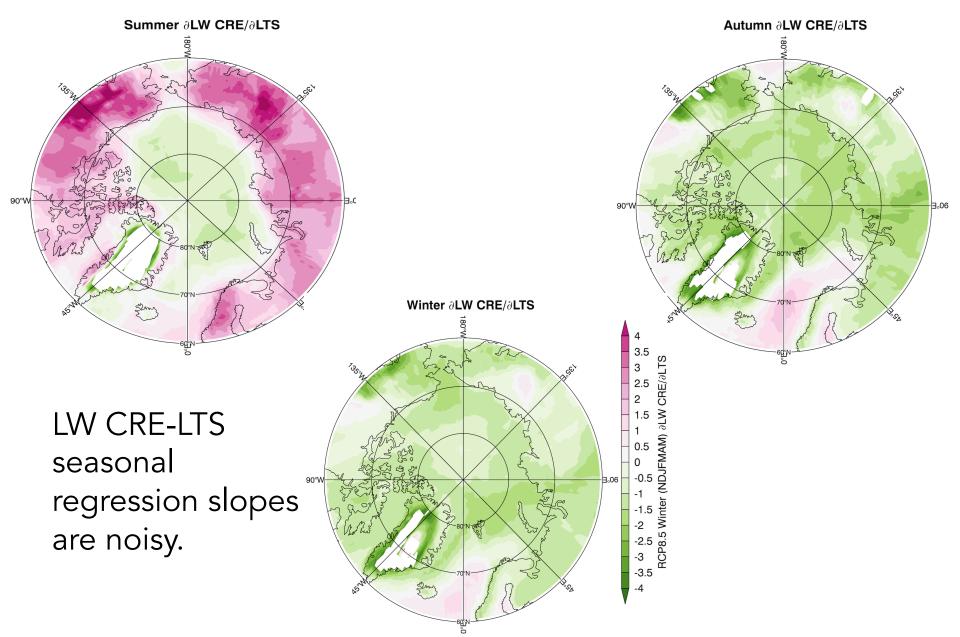


The patterns from the ensemble average and regression result agree well and the regression residuals are ~10-40%.

Inter-annual variability of LTS and LW CRE:

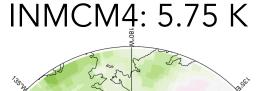


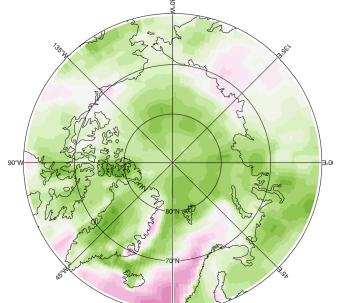
Inter-annual variability of LTS and LW CRE: CMIP5 RCP8.5



Inter-annual variability of LTS and LW CRE:

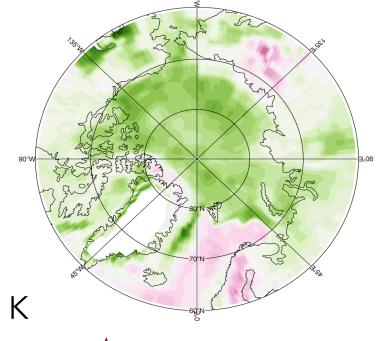
CMIP5 RCP8.5



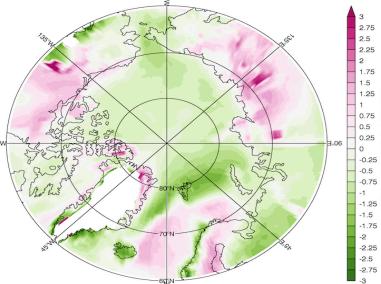


Model-to-model variations are largest over the Central Arctic and in the Barents Sea MIZ





CCSM4: 7.10 K



Takeaway messages:

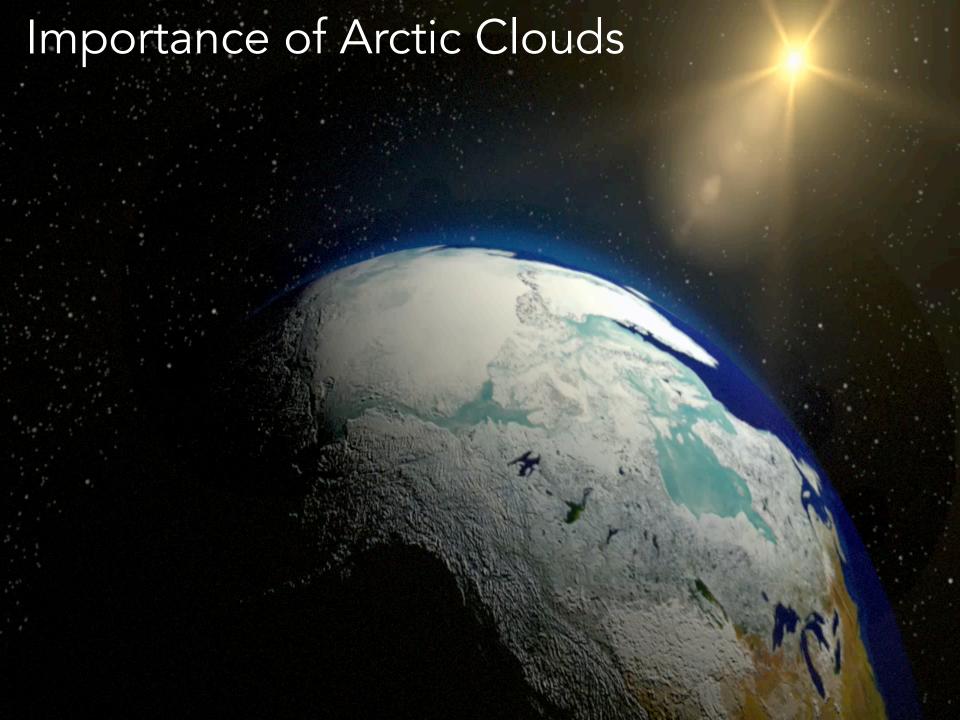
- Lower tropospheric stability, an excellent diagnostic for Arctic low cloud properties and surface radiative effects.
- Small change in atmospheric states can equal a large cloud feedback! A small change in the frequency of occurrence of atmosphere states can yield a larger Arctic cloud feedback than any cloud response to sea ice.
- Too strong LW CRE response to a reductions in Arctic stability.
 Over the central Arctic, the LW CRE sensitivity to a change in
 LTS in climate models is much stronger (more than 2 times in
 places) than observations show.
- Potentially significant amplifying sea ice-cloud feedback in fall.
 Our results indicate that the cloud response to sea ice
 reduction in fall could delay the fall freeze-up and influence the
 variability in sea ice extent and volume, under certain
 meteorological conditions; It only takes 1 Wm⁻² to make a big
 difference.

Takeaway messages:

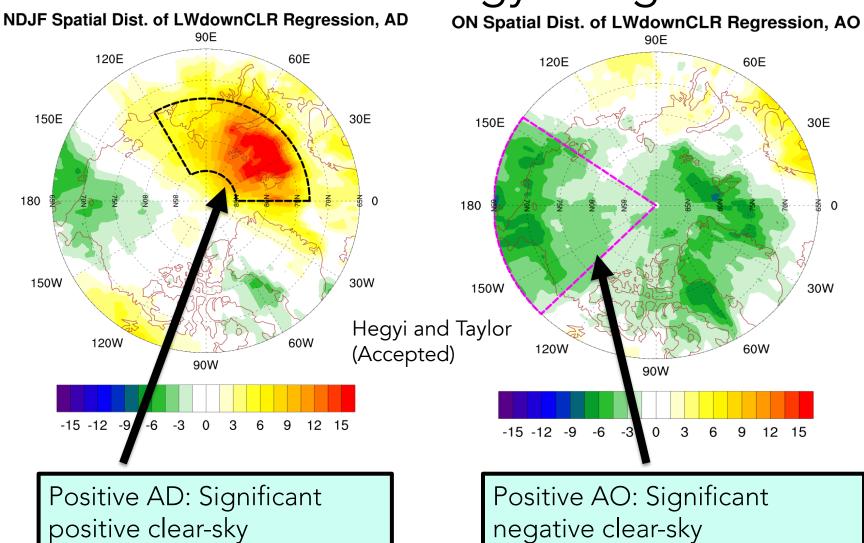
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Any way the wind blows: Circulation and the Arctic Surface Energy Budget

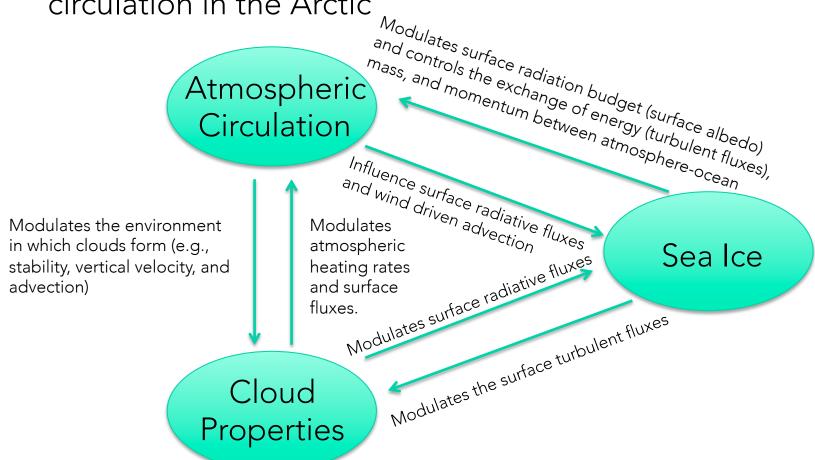


downwelling LW anomaly.

downwelling LW anomaly.

How do we reach the finish line?

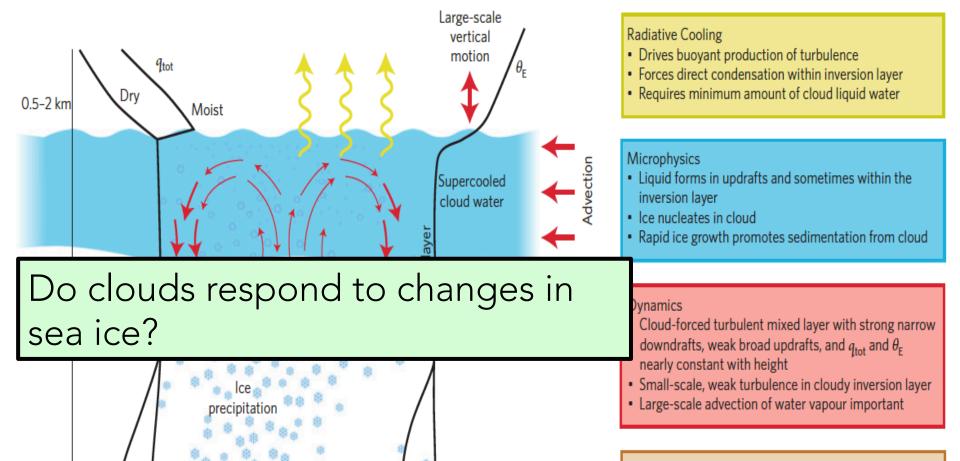
Understanding the coupling between the cloud and circulation in the Arctic



In the Arctic, we cannot consider the interactions between clouds and the circulation without considering the sea ice state because sea ice influences both clouds and circulations.

Arctic Low Cloud Processes

Deconpled



onpled

The influence of the surface type on the cloud properties implies an interaction between clouds and sea ice that may significantly influence Arctic climate change.

Morrison et al. (2012; Nature Geoscience)

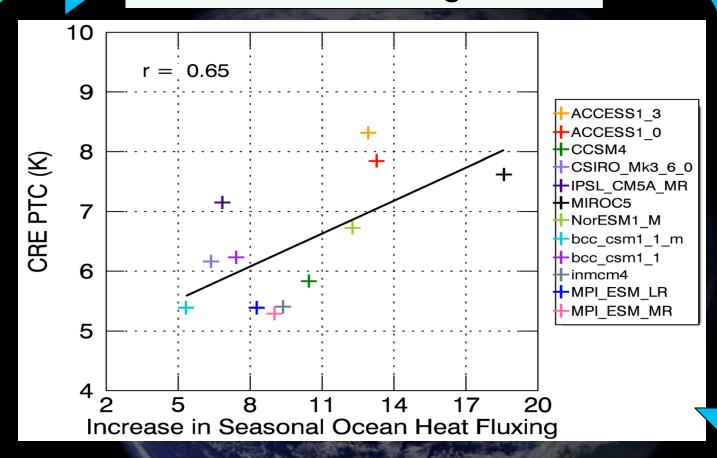
Turbulence and q contributions can be weak or strong

Sink of atmospheric moisture due to ice precipitation
Surface type (ocean, ice, land) influences interaction

Surface Layer

with cloud

Larger CRE results in more summer storage



Increases in clouds strengthen the CRE PTC

Increases in storage facilitates cloud formation and IR feedbacks

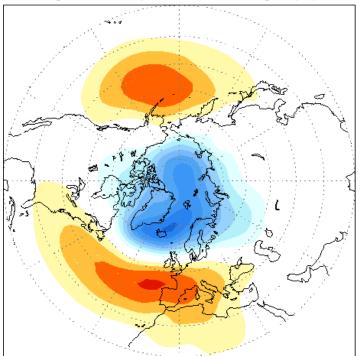


Hegyi and Tayor (GRL; accepted)

Defining the Arctic circulation

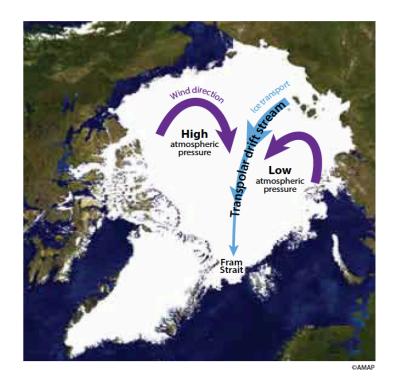
AO: Leading EOF of the 1000 hPa height pattern from 20-90N.

Leading EOF (19%) shown as regression map of 1000mb height (m)



Zonally symmetric

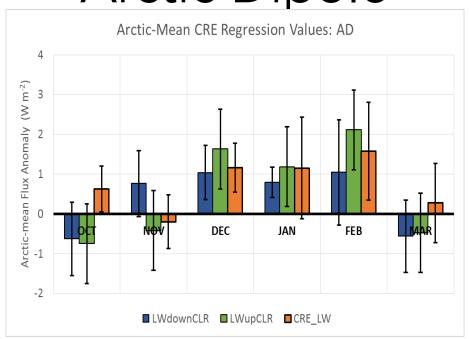
AD: Leading EOF of the 1000 hPa height pattern from 70-90N.



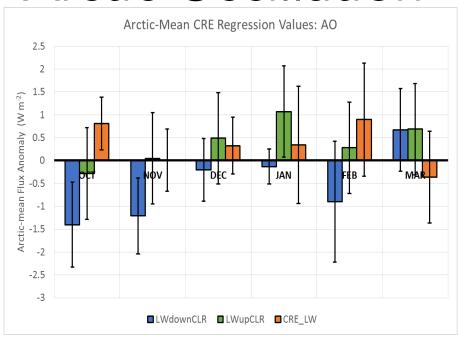
Zonally asymmetric

Any way the wind blows: Circulation and the Arctic Surface Energy Budget

Arctic Dipole

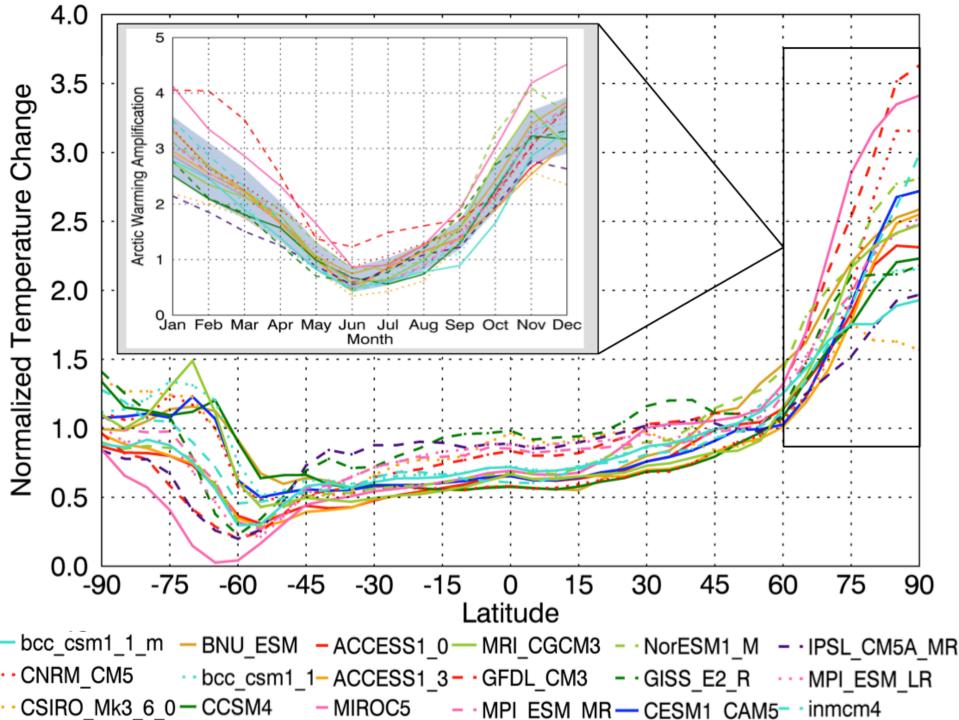


Arctic Oscillation



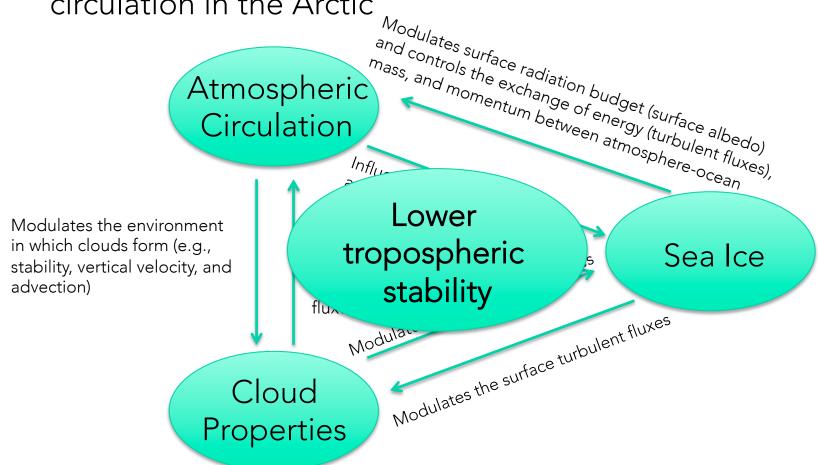
Robust, domain averaged associations between the Arctic circulation and the surface radiation budget.

Hegyi and Taylor (GRL; accepted)



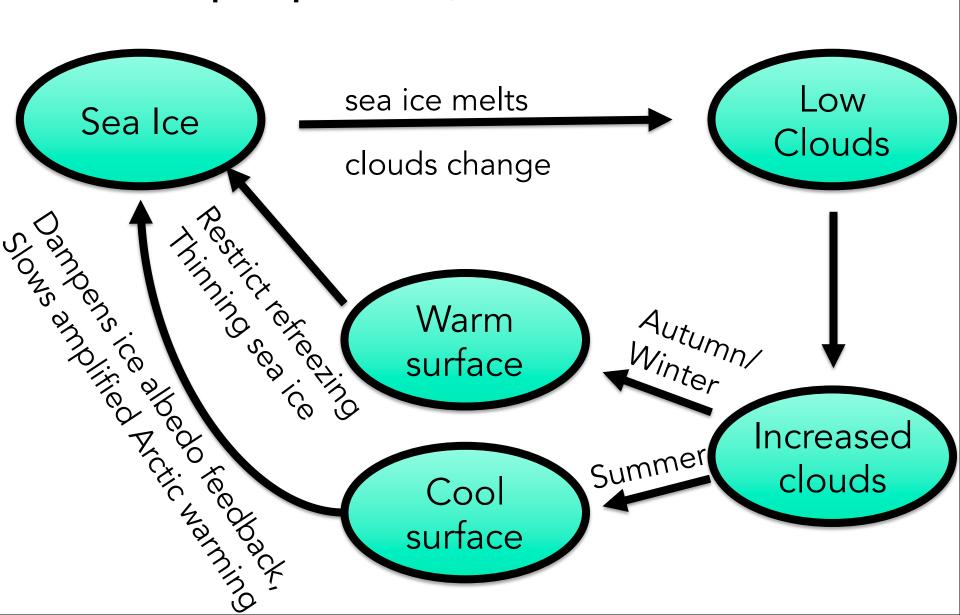
Another 3-legged stool?

Understanding the coupling between the cloud and circulation in the Arctic

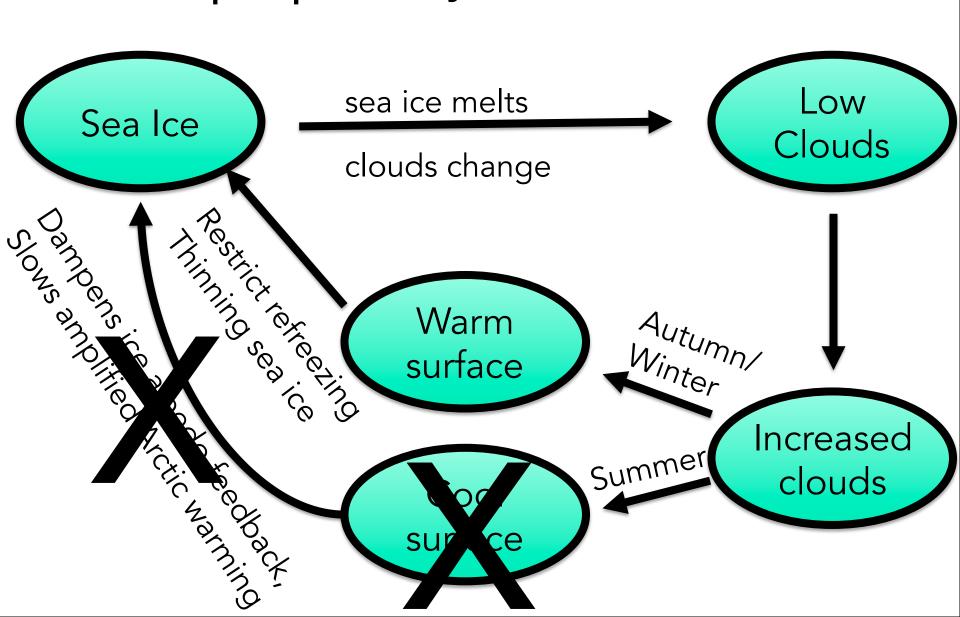


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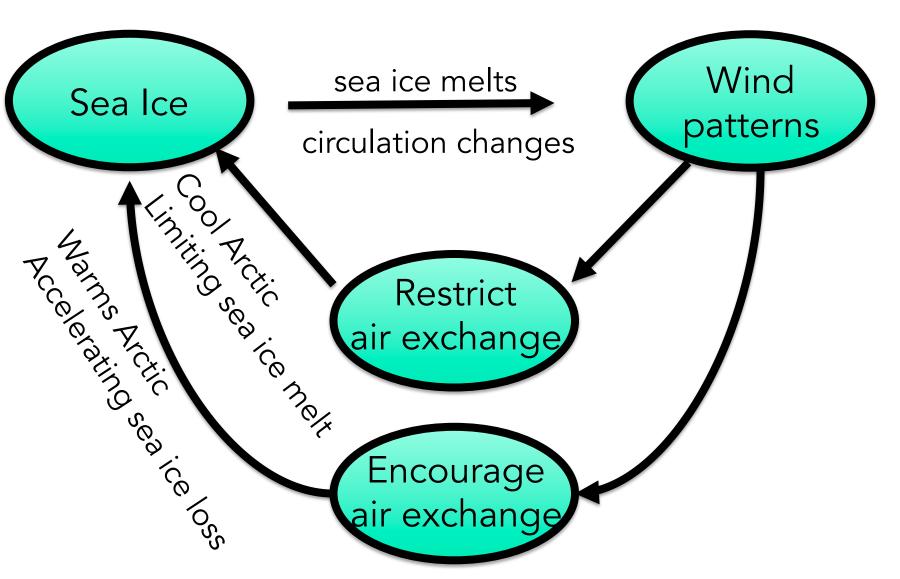
Example pathways: Sea Ice-Clouds



Example pathways: Sea Ice-Clouds



Example pathways: Sea Ice-Circulation



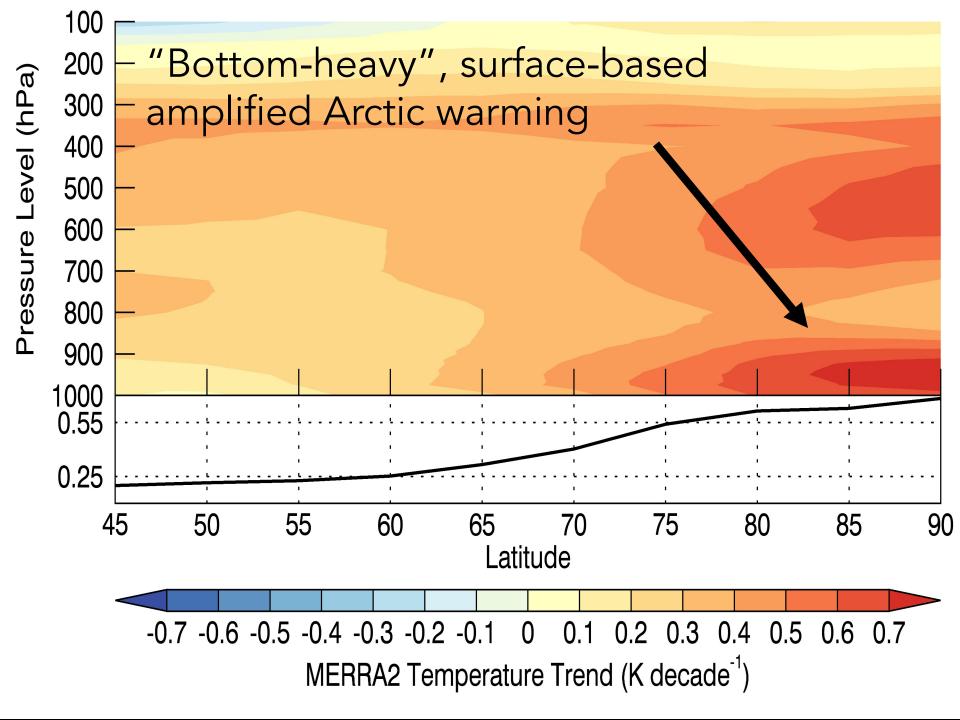
CLOUD AMOUNTS DECREASE

CLOUD WATER CONTENT DECREASES

CLOUD HEIGHT DECREASES

Taylor et al. (2015)

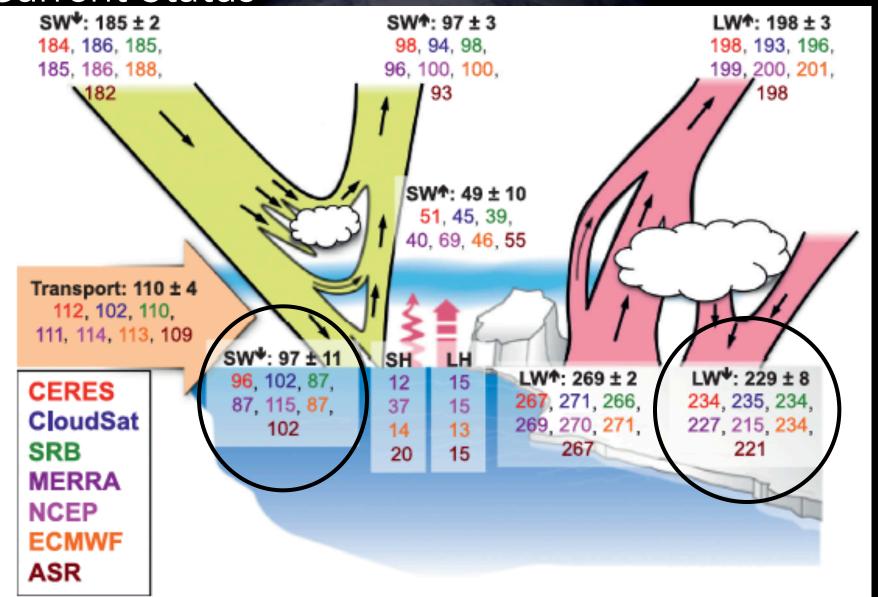
SEA ICE INCREASES



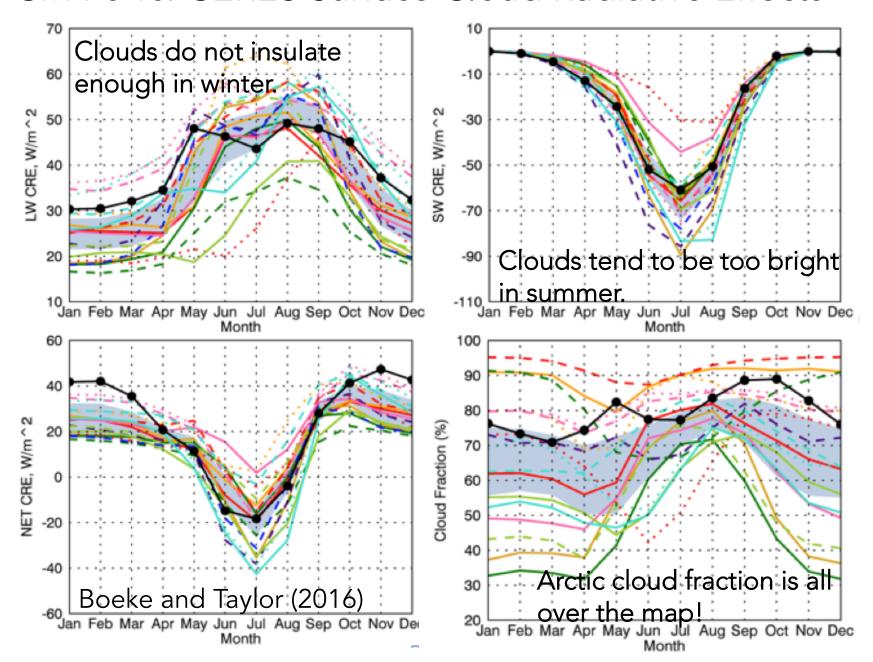
Arctic surface energy budget:

Current Status

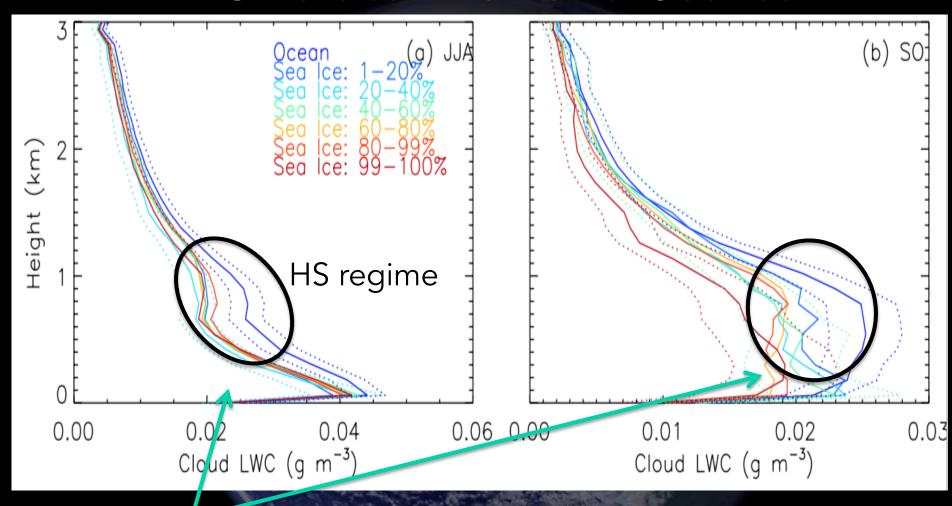
Christensen et al. (2016; BAMS)



CMIP5 vs. CERES Surface Cloud Radiative Effects

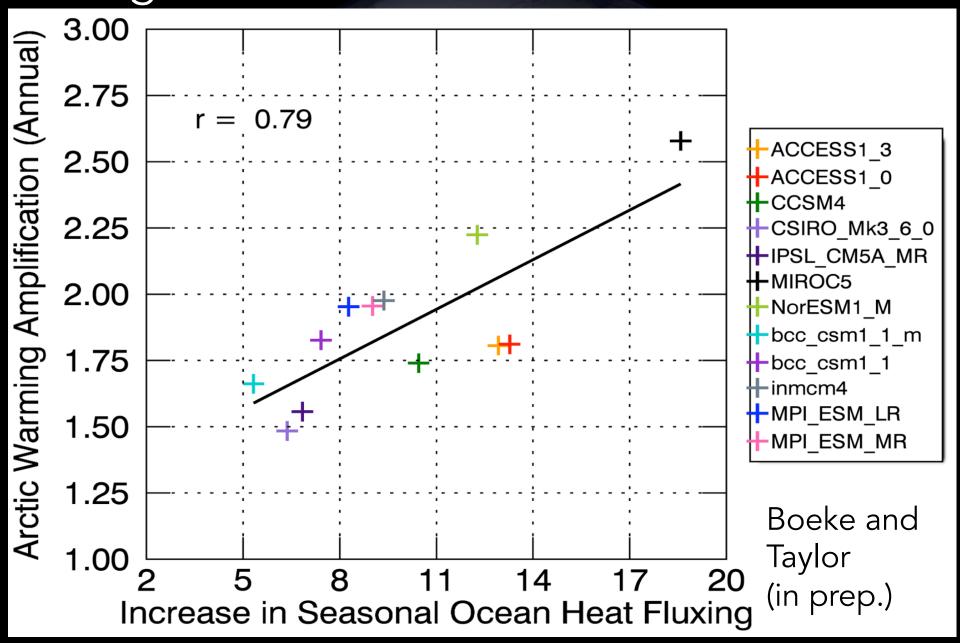


LWC Vertical Profile vs. Sea ice



- General decrease in LWC is found with increased sea ice concentration in both summer and autumn.
- Statistically significant differences the LWC between 500 m and 1.2 km are found in summer and autumn at 0% and 20-40% sea ice concentration.

A big role for the ocean!



Relationship between LTS-CRE by atmospheric regime

Units: W m⁻²

Atmospheric	Summer		Fall (SO)		Winter	
Regime	(JJA)				(NDJFMAM)	
	LW	SW	LW	SW	LW	SW
	CRE	CRE	CRE	CRE	CRE	CRE
VHS (LTS >24 K)	40.6	-57.5	37.9	-10.8	24.4	-8.1
HS	50.1	-70.1	53.0	-16.6	37.9	-11.8
(16 <lts<24 k)<="" th=""><th></th><th></th><th></th><th></th><th></th><th></th></lts<24>						
S (LTS<16 K)	55.7	-80.8	59.5	-20.2	53.6	-19.7
UL (ω_{500} < -8.6	51.8	-85.1	56.6	-22.3	48.6	-18.6
hPa day ⁻¹)						